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Fossil mantle-sediments interface recognized in the Western Alps metaophiolites: a key to unravel the accretion mechanism of the Jurassic Tethys ocean

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ABSTRACT

In the southern Aosta Valley (Italian Northwestern Alps), metaophiolites are mainly composed of serpentinitized mantle-derived peridotites intruded by gabbros and rodingitic dykes, well exposed in the Mount Avic area, and of smaller amounts of mafic rocks and metatrandhjemite. This rock assemblage recalls the “slow-spreading” lithosphere created at modern mid-ocean ridges. Meta-ophiolites show a dominant early Alpine subduction-related metamorphic imprint under eclogite/blueschist facies conditions, variously retrogressed under greenschists facies conditions. In the high Champorcher Valley (SW of Mount Avic) serpentinites are directly covered by a serpentinite *mélange* followed by flysch-like calcschists with detrital ophiolitic interbeds. Despite the pervasive Alpine tectonic deformation and metamorphic recrystallization through subduction-related stretching and boudinage and collision-related folding, the *mélange* internal fabric still retains records of a block-in-matrix structure, well consistent with mass-transport processes related to an active oceanic tectonic setting in which mantle rocks were progressively and continuously exhumed by faulting. The products of mass-transport processes and faulting are unconformably sealed by flysch-type calcschists embedding cm-sized clasts of actinolite/tremolite-schists interpreted as detrital ophiolitic material. The serpentinite *mélange* is interpreted as syn-extensional sedimentary rocks produced at the mantle-sediments interface on the Jurassic Tethys ocean floor and subsequently overprinted by subduction zone tectonics.

KEY WORDS: meta-ophiolite, serpentinite *mélange*, Western Alps.

INTRODUCTION

The Western Alps meta-ophiolites have been emplaced during the closure of the *Ligure-Piemontese* basin, a segment of the Jurassic Tethys ocean interposed between the European and African plates. Since 1980s to early 1990s, after many decades of investigations conducted on the Alpine meta-ophiolites, and particularly since the advances in marine geology techniques favouring a better knowledge of modern oceanic crust, the scientific community has become increasingly aware of the fact that the Tethyan ocean lithosphere did not fit the conceptual model of a “normal” or layered oceanic crust, as defined by the Penrose Conference (Anonymous, 1972; Lagabriele & Cannat, 1990). Instead, Tethyan meta-ophiolites are characterized by dominant mantle-derived (mostly serpentinitized) peridotites associated with

gabbro intrusions, and by less amounts of mafic rocks deriving from lava flows, pillow lavas, and gabbros. Such internal architecture recalls the structure and composition of modern “magma-poor” (e.g., Dick et al., 1981; Mével et al., 1991) or “slow-spreading ridge” (e.g., Cannat, 1993) oceanic crust.

Metasediments overlying meta-ophiolites consist of metaquartzites, marble, and calcschists (with various proportions) which directly cover metabasalts or, alternatively, mantle rocks and gabbros. Field evidence for mantle rocks exposition on the Tethyan seafloor are constrained by the occurrence of ophicalcites and ophicarbonates breccias interposed between the ophiolitic basement and the metasedimentary cover (e.g., Driesner, 1993; Tartarotti et al., 1998; Dal Piaz, 1999, for the northwestern Alps). In present-day oceans, mantle peridotites and gabbros have been recovered within a variety of tectonic settings including transform faults, the walls of mid ocean ridges, ridge-transform intersections (e.g., Karson et al., 1987; Mével et al., 1991; Bonatti et al., 1990; Blackman et al., 1998). Exposure of mantle rocks on the seafloor and mantle denudation mechanism have important tectonic and geochemical implications for lithospheric extension and subduction process (e.g., Lemoine et al., 1987; Deschamps et al., 2013; Schwartz et al., 2001). Therefore, informations from modern oceans are fundamental for understanding the geodynamic evolution of the Tethyan lithosphere.

Here we focus on a portion of the Western Alps meta-ophiolites where serpentinitized mantle peridotites are directly covered by ophicarbonates breccias and calcschists. On the basis of lithostratigraphic and petrographic features this contact is interpreted as a fossil basement-cover interface. Inferences on the paleo-topography and physiography of the Tethys ocean seafloor are then discussed.

REGIONAL GEOLOGY

The study area is located in the southern Aosta Valley (Northwestern Alps, to the south of the Aosta-Ranzola fault system) and comprises metamorphic ophiolites comparable to the Zermatt-Saas (ZS) unit (e.g., Bearth, 1967; Dal Piaz & Ernst, 1978). The ZS unit consists of antigorite-Ti-clinohumite-

magnetite-bearing serpentinites, Fe-Ti- and Mg-metagabbros, metabasalts, covered by metasedimentary rocks showing a dominant early Alpine metamorphic imprint under eclogite-blueschist-facies conditions, partially retrogressed under greenschists facies conditions.

Meta-ophiolites exposed in the southern Aosta Valley are mainly composed of serpentinitized mantle-derived peridotites intruded by gabbros and rodingitic dykes, well exposed in the Mount Avic area (Fontana et al., 2008; Fontana et al., 2015), and of smaller amounts of mafic rocks and metatrandhjemite (Dal Piaz et al., 2010; Novo et al., 1989). These meta-ophiolites show a dominant early Alpine subduction-related metamorphic imprint under eclogite-blueschist-facies conditions of Eocene age, variously retrogressed under greenschists facies conditions (e.g., Baldelli et al., 1985; Benciolini et al., 1988; Martin & Tartarotti, 1989; Dal Piaz et al., 2001).

Two main types of metasedimentary successions are recognizable in the southern Aosta Valley: one succession comprises metaradiolarite, marble, calcschists s.s., and flysch-like calcschists. This suite generally occurs on top of mafic rocks, as in the St. Marcel Valley (Martin & Tartarotti, 1989; Tartarotti et al., 1986, 2014; Tartarotti & Caucia, 1993), where metabasalts and metasediments are also associated with sulphide- and Mn-rich ore deposits (Martin et al., 2008; Martin & Tartarotti, 1989; Rebay & Powell, 2012; Tartarotti & Caucia, 1993; Tumiatì et al., 2005, 2010). This ophiolitic sequence likely refers to the shallowest part of the Tethyan oceanic crust created near the ridge axis where hydrothermal fluid circulation was active and covered by pelagic sediments. The other type of metasedimentary succession consists of ophicarbonates breccias which directly lie upon serpentinitized mantle peridotites, followed by flysch-like calcschists with ophiolitic interbeds (Tartarotti et al., 2015). We here describe the second type of metasedimentary succession which occurs in the high Champorcher valley (SW of the Mount Avic serpentinite massif), in the Lac Miserin area.

THE LAC MISERIN SERPENTINITIC MÉLANGE

In the Champorcher Valley, meta-ophiolites are made of metabasalt, metagabbro, and serpentinite (Dal Piaz & Nervo, 1971; Nervo & Polino, 1976; Dal Piaz et al., 1979). Serpentinites are extensively exposed in the huge Mount Avic massif (on the left side of the Champorcher valley) and in small outcrops in the Lac Miserin area (right side of the high Champorcher Valley). These latter are interpreted as tectonic slices transposed from the Mount Avic serpentinite. In the Lac Miserin area, serpentinites are directly covered by a serpentinite mélangé followed by flysch-like calcschists with detrital ophiolitic interbeds. Upon massive to veined and brecciated serpentinite, sealed by metasomatic rocks rich in tremolite and carbonates, the mélangé starts with white marble and carbonatic calcschist alternating with decimeters-thick horizons of serpentinite metabreccia and metagraywacke, including isolated metric blocks of serpentinite (Figure 1). The sequence continues with rounded to irregularly-shaped blocks of serpentinite, cm-to several dm in size, randomly distributed within a matrix of foliated impure marbles and calcschist. It is then unconformably covered by flysch-like calcschists

embedding cm-sized clasts of actinolite/tremolite-schists, interpreted as detrital ophiolitic material.

The serpentinite mélangé has been severely overprinted by Alpine tectonic deformation and metamorphic recrystallization through subduction-related stretching and boudinage and collision-related folding. At least four Alpine deformation phases were recognized in the field which produces isoclinal folds followed by asymmetric folding overprinted by wide open folds and late S-C structures. Rootless and isoclinal folds constrain the timing of the serpentinite mélangé emplacement to the pre-Alpine or early Alpine history. Despite this intense tectonic reworking, the mélangé internal fabric still retains records of breccia and block-in-matrix structures, well consistent with mass-transport processes related to an active tectonic setting in which mantle rocks were progressively and continuously exhumed by faulting.



Fig. 1 - Outcrop of the Lac Miserin serpentinite mélangé showing serpentinite metabreccia and metagraywacke and serpentinite blocks (Foto P. Tartarotti)

In conclusion, the serpentinite/calcschists juxtaposition in the Lac Miserin area is separated by an interposed serpentinite mélangé which thus corresponds to a fossil mantle-sediments interface established in the Jurassic Tethys ocean. We infer that the serpentinite mélangé corresponds to syn-extensional sedimentary rocks produced at the mantle-sediments interface on the Jurassic Tethys ocean floor and subsequently overprinted by subduction zone tectonics.

This scenario has profound implications for the physiography of the Tethys Ocean. Our results indicate that the Tethyan seafloor should have been characterized by extensive regions of active faulting and mass wasting responsible for the formation of a serpentinitic rugged seafloor. Serpentinite metabreccia and metagraywacke of the Lac Miserin mélangé may represent talus breccias accumulated at the base of fault scarps, then indurated by carbonate sediment (ooze?), whereas larger blocks of serpentinite randomly distributed within foliated impure marbles and calcschist may be interpreted as coarse, poorly sorted blocks deposited by mass wasting. Similar deposits have been extensively observed and sampled in the Atlantic Ocean (e.g., Karson & Lawrence, 1997). Ophicarbonates rocks resting upon the serpentinite basement indicate that the upper mantle peridotites were already

exhumed on the seafloor prior to the onset of subduction zone tectonics within the ocean basin. Finally, flysch-like calcschists are here interpreted as a post-extensional sedimentary sequence sealing the serpentinite-mélange interface.

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